

MET3003



Lecture 11: Hydrometeors
21FEB07

Atmospheric Aerosols

- Suspensions of solid or liquid in a gas, in this case air.
 - Clouds are aerosols, but in meteorology we usually mean different chemical composition than H_2O .
- **Sizes:** measured in μm ($= 10^{-3}$ mm) to $0.01 \mu m$. (human red blood cell $\sim 5 \mu m$).
- Because of friction and low mass aerosol particles fall 0.03 to 3 mm s^{-1} (i.e. a few meters a day)
- **Concentration:**
 - Oceans 10^3 cm^{-3}
 - Inland (clean) 10^4 cm^{-3}
 - Inland (polluted) 10^5 cm^{-3}
 - Primarily below 3-5 km altitude
- **Sources:** combustion, sea salt, some soils, decomposition of some plants
- **Sinks:** Coagulation, capture in clouds, fallout, precipitation scavenging

Some Definitions

- **Hydrometeor:** Any product of condensation, deposition, or freezing from the atmosphere whether in the atmosphere or on the Earth's surface. Also any water substance blown from the Earth's surface by wind---e.g. windblown spray or snow.
- **Precipitation:** Hydrometeors that fall from a cloud and reach the ground.
- Visible clouds are made of:
 - **Cloud water:** Suspended droplets $10 \mu m$ in diameter (Middle Latitudes) or $15\text{--}60 \mu m$ in diameter (Tropics) that **condensed** directly from vapor
 - **Cloud ice:** Comparably sized crystals that **deposited** directly from vapor
 - Unfrozen liquid (**Supercooled**) water at temperatures below $0^\circ C$ is common, particularly outside the tropics

Hydrometeors that are NOT Precipitation

- **Dew:** Liquid H_2O that condenses directly on the ground.
- **Frost:** Ice that is deposited directly on the ground from vapor
- **Rime:** Supercooled fog that freezes on contact with the surface
- **Black Frost** or **Frozen Dew:** Condenses then freezes.

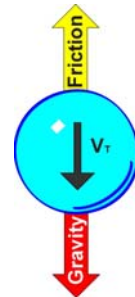


Kinds of Precipitation

- **Rain:** Liquid water drops $> 0.5 \text{ mm}$ in diameter falling from the sky
- **Drizzle:** Liquid water drops $< 0.5 \text{ mm}$ in diameter falling from the sky
- **Virga:** Like rain but evaporates before it reaches the ground.
 - Virga is a hydrometeor, but it isn't precipitation.
- **Hail:** More or less spherical ice $> 5 \text{ mm}$ in diameter
- **Graupel:** More or less spherical ice $< 5 \text{ mm}$ in diameter
Note that the boundary between Hail and Graupel is 10x larger than the boundary between Rain and Drizzle
- **Snow:** Ice crystal deposited directly from vapor, or aggregates (clumps) of crystals
- Note that the boundary between Hail and Graupel is 10 times larger than the boundary between Rain and Drizzle

Terminal Velocity

- Hydrometeors fall speed is a balance between friction and gravity.
 - Cloud particles: mm or cm s^{-1}
 - Rain up to 8 m s^{-1}
 - Snow $1\text{--}2 \text{ m s}^{-1}$
 - Hail $> 10 \text{ m s}^{-1}$
- VT increases for
 - Lower air density
 - Hydrometeor size
 - Smoother hydrometeors



What Happens in a Cloud?

- Updraft leads to cooling and condensation.
- Forms cloud droplets, 10-50 μm in diameter.
- Cloud droplets collide and coalesce to form rain
- Are also carried upward to temperatures $< 0^\circ\text{C}$
- Many eventually freeze
- Ice grows faster than supercooled liquid H_2O
- Forms snow, graupel, and hail
- Which may melt to form rain if it falls below 0°C level
- Precipitation loading and evaporative cooling force downdrafts.



Two Kinds of Precipitation

- Stratiform
 - Forced by large-scale flow
 - Updrafts a few cm s^{-1}
 - Falls from Nimbostratus
 - Drizzle, Rain & Snow
 - Cool season (sort-of)
- Convective
 - Buoyant updrafts
 - Updrafts several m s^{-1}
 - Falls from cumulonimbus
 - Rain, graupel & hail
 - Warm season

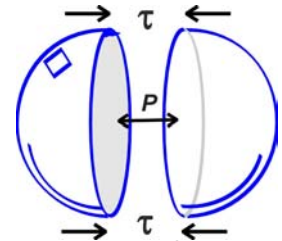


What are clouds made of?

- **Not of raindrops**, but of **suspended cloud droplets**---about 100 times smaller than raindrops, i.e. 10-80 μm (0.01-0.08 mm) in diameter.
- Let's start with cloud droplet formation in supersaturated air through **Condensation of water vapor**:
- **Homogeneous nucleation**: Water drops (or ice crystals) *might* form in pure air with nothing but dry air and H_2O involved.
- Each water molecule added has to do work against the surface tension.
- Thus, it is hard to make cloud droplets from scratch, because the amount of work per molecule is more when the droplet is smaller.
- Homogeneous nucleation requires **> 110% relative humidity** (or 10% supersaturation), which generally does not occur in the atmosphere.

It takes a lot of work against surface tension to add water to a very small drop

Very small droplets have so much surface area relative to their volumes and such high internal pressure that more H_2O atoms diffuse outward than inward.



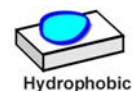
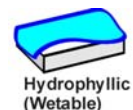
$$2\pi r\tau = \pi r^2 p, \quad \text{work} = p dV = \frac{2\tau}{r} dV$$

Heterogeneous nucleation

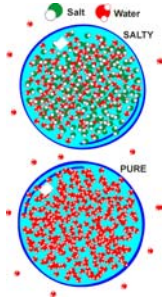
- Water drops (or ice crystals) form from the vapor on some sort of **substrate** or by **soaking** into a salt particle.
- For **wetable** nuclei, water attaches to the surface.
- For **hygroscopic** nuclei, water soaks into the surface and may actually dissolve the nucleus.
- What makes a good **Cloud Condensation Nucleus (CCN)**?
 - Reasonably big (for an aerosol particle)
 - Easily wettable (0.1 μm) or soluble (0.01 μm).
- Over land, 100-1000 CCN cm^{-3} typically become active at 1% supersaturation
- Over the oceans, $< 100 \text{ CCN cm}^{-3}$ typically become active at 1% supersaturation

Kinds of Condensation Nuclei

- For **Hydrophillic (wetable)** nuclei, water attaches to the surface.
- For **hydrophobic** nuclei, water may bead on the surface,
 - But doesn't attach
 - Not good condensation nuclei
- For **hygroscopic** nuclei, water soaks into the surface and eventually dissolves the nucleus.



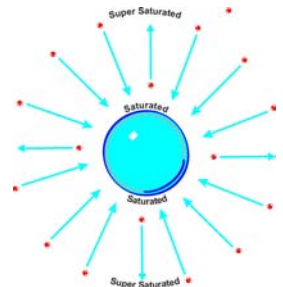
Solute effect:



- Dissolved "salt" blocks part of the surface, preventing H_2O molecules from escaping. Effective saturation vapor pressure is reduced.
- Proportional to: (number of H_2O molecules)/(number of H_2O + salt molecules) =
- Effect is weaker as the droplet grows, diluting the salt

Growth of Cloud Droplets by Diffusion

- The cloud becomes super-saturated
- Water vapor diffuses from high vapor density to low vapor density
- Rate of mass growth is proportional
 - Droplet area
 - Supersaturation
- Rate of growth of radius
 - Area is proportional to radius squared
 - Volume is proportional to radius cubed
 - Radius growth is inversely proportional to radius itself.
 - It's a bit trickier than the foregoing because the rate of diffusion also becomes smaller for larger droplets

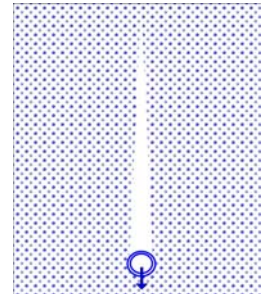


Limits to Growth by Diffusion

- The slowing of the growth rate means that there is a limit to the size drops can reach by diffusion.
- **NOTE: They can't form at all if they start out too small and they can't grow too big**
- **Continental clouds:**
 - 100-1000 CCN cm^{-3} ,
 - Small supersaturations,
 - All of the droplets are about 10 μm diameter
- **Maritime clouds:**
 - <100 CCN cm^{-3}
 - slightly larger supersaturations,
 - Wide range of droplet diameters, 20-50 μm .
- Both typically contain 1-2 $gm\ m^{-3}$ of liquid water
- The biggest of these falls only about 30 $cm\ s^{-1}$, and cloud updrafts are often several $m\ s^{-1}$.
- So how do we get rain?

Growth of Warm Rain

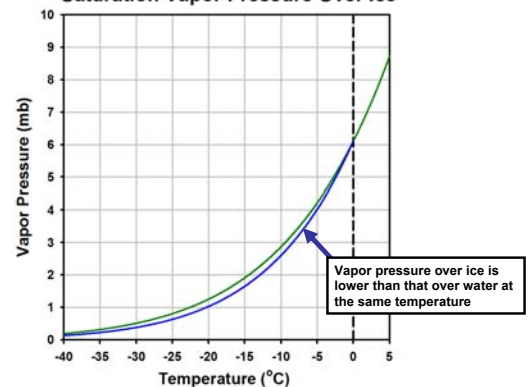
- **Warm Rain:** Falls from clouds where $T > 0^\circ C$ throughout
- **Autoconversion:** A few droplets collide randomly and merge, becoming larger than the others
- Hydrometeors fall near their terminal velocity
- Larger, denser or rounder hydrometeors fall faster
- In clouds warmer than $0^\circ C$, precipitation forms when larger particles sweep up smaller ones as they fall through the cloud.
- Called **Collision Coalescence**.



Collision-Coalescence

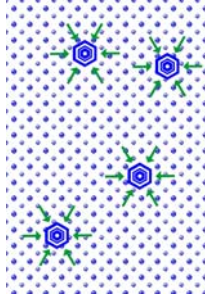
- Droplets bump into each other and merge to make bigger drops:
- Terminal velocity (6.5 $m\ s^{-1}$ for a 1mm rain drop) of the larger drop allows it to overtake and capture cloud droplets.
- Some, or all, of the terminal velocities may be less than the updrafts velocity so that all of this can be moving upward, at least until the drops get big.
- **Coalescence efficiency** = fraction of drops or droplets in the cone that actually merge with collector drop.
- Varies from 0.2 for small droplets to > 1 for drops that are nearly the same size.
- Really big drops (> 2 mm) break up, and pieces are sometimes carried back up into the cloud for another go.
- Because of the broader initial droplet distribution, C-C works with a vengeance in tropical cumuli.

Saturation Vapor Pressure Over Ice



Mixed-Phase, Clouds, $T < 0^{\circ}\text{C}$

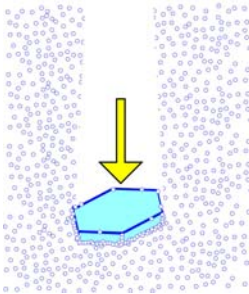
- Both freezing and condensation generally require nuclei to start the process
- Natural clouds have plenty of condensation nuclei (100,000/liter) but few freezing nuclei (10-100/liter)
- Consequently many clouds colder than 0°C contain **Supercooled** liquid H_2O that has not yet frozen.
- May be as cold as -20°C
- When some ice does form,
 - the ice crystals grow
 - at the expense of supercooled water droplets
 - Bergeron-Findeisen process
- Artificial freezing nuclei can promote growth of snow that melts to form rain



Precipitation Growth Processes

- Autoconversion:** When cloud droplets or ice crystals clump to form larger cloud droplets or drizzle
- Collision-coalescence:** Rain drops grow by collecting cloud droplets or smaller rain drops
- Riming:** Hail or graupel collect supercooled water that freezes on the collecting hydrometeor
- Accretion:** Frozen hydrometeors collect cloud ice crystals or smaller frozen hydrometeors

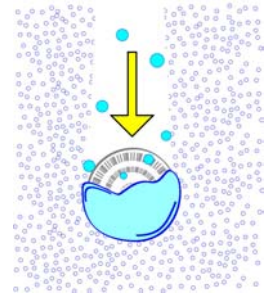
Riming



- Starts with a vapor-grown ice crystal
- Falling through a supercooled cloud
- Droplets freeze on leading edges.
- Forming spongy ice
- This is how graupel forms.

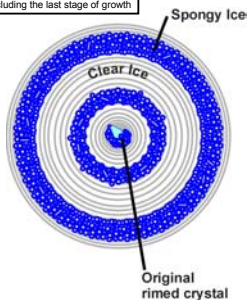
Wet Growth or Wet Riming

- Temperatures $> -5^{\circ}\text{C}$
- Latent heat released by freezing keeps hailstone at 0°C .
- Coated with liquid water
 - Some freezes
 - Some is shed as drops at 0°C



Private Lives of Hailstones

Two episodes of "dry" riming plus initial growth
Three episodes of wet riming including the last stage of growth



- Updrafts in Cb's are often stronger than hydrometeor terminal velocities.
- Leads to recycling of graupel and hail
- Forming multiple layers of clear and spongy ice
- In very strong updrafts hailstones can get really big---
- Coffeyville KS, 22JUN03
 - 18 cm diameter,
 - 750 gm

Hail Images

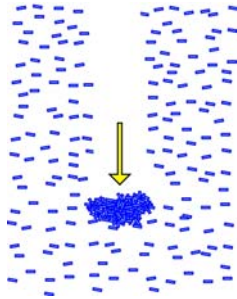


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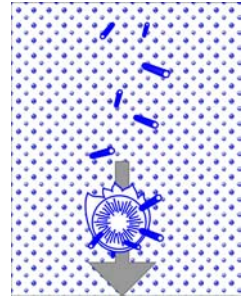
Aggregation

- Falling frozen hydrometeor collects other frozen hydrometeors
- Most snowflakes are aggregates of vapor-grown crystals
- Works best at temperatures $> -5^{\circ}\text{C}$
- Big, soggy flakes vs. diamond dust

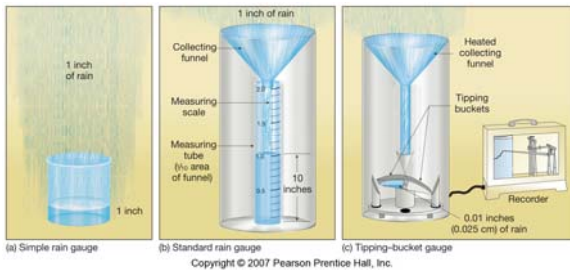


Ice Multiplication

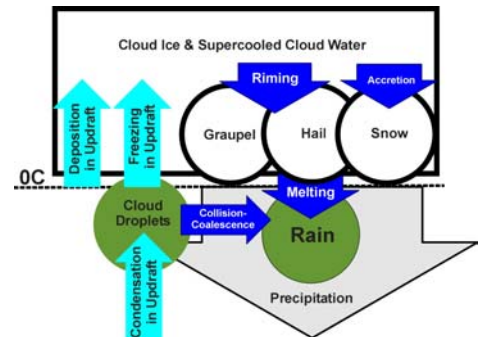
- Observations showed little ice in tropical Cb's
- Hallett-Mossup Process
 - Graupel falling in supercooled clouds grow whiskers-like ice crystals
 - Break off
 - Natural cloud seeding
- Happens in maritime clouds only
- Doesn't look good for cloud seeding in the tropics



Rain Gauges



Precipitation Formation



Summary

- Know the definitions of **Aerosol**, **Precipitation** and **Hydrometeors**
- Know the definitions of **Frost**, **Dew**, **Virga**, **Cloud Water (droplets)** **Cloud Ice (crystals)** and **Terminal Velocity**.
- Be able to explain **Supercooled Water**, **Homogeneous** and **Heterogeneous Nucleation**.
- Be able to explain how the different structure of maritime and continental cumuli arise from CCN concentrations in the parent air mass.
- Understand the growth of hydrometeors from vapor
- Be able to explain the **Bergeron-Findeisen Process** (ice crystal)
- Be able to explain the difference between **Stratiform** and **Convective** precipitation.
- Be able to explain hydrometeor growth by **Autoconversion**, **Collision-Coalescence**, **Accretion**, and **Riming**
- Be able to interpret the history of a hailstone from a cross section showing layers of clear and spongy ice
- Be able to explain how precipitation is measured with a rain gauge.
- Reading for next time: L&T Chapter 6, Winds and Pressure**